

1 **Determination of the $\delta^2\text{H}$ values of high molecular weight lipids by high temperature**
2 **GC coupled to isotope ratio mass spectrometry**

3 Sabine K. Lengger^{1,2*}, Yuki Weber³, Kyle W.R. Taylor⁴, Sebastian H. Kopf⁵, Robert Berstan⁴,
4 Ian D. Bull¹, Jan-Peter Mayser¹, William D. Leavitt⁶, Jerome Blewett¹, Ann Pearson³ and
5 Richard D. Pancost^{1,7}

6 1 Organic Geochemistry Unit, School of Chemistry, University of Bristol, Cantock's Close,
7 Bristol BS81TS, UK

8 2 Biogeochemistry Research Centre, School of Geography, Earth and Environmental
9 Science, University of Plymouth, Drake Circus, Plymouth PL48AA, UK

10 3 Department of Earth and Planetary Sciences, Harvard University, 20 Oxford St,
11 Cambridge, MA 02138, USA

12 4 Elementar UK Ltd., Earl Road, Cheadle Hulme, Stockport, SK8 6PT, UK

13 5 Department of Geological Sciences, University of Colorado Boulder, Boulder, CO, USA

14 6 Department of Earth Science, Department of Chemistry, Department of Biological
15 Sciences, Dartmouth College, Hanover, NH, USA

16 7 School of Earth Sciences and Cabot Institute for the Environment, University of Bristol,
17 Queens Road, Bristol BS81RL, UK

18 * corresponding author: sabine.lengger@plymouth.ac.uk

19

20 **Abstract**

21 Rationale: The hydrogen isotopic composition of lipids ($\delta^2\text{H}_{\text{lipid}}$) is widely used in food
22 science and as a proxy for past hydrological conditions. Determining the $\delta^2\text{H}$ values of large,
23 well-preserved triacylglycerides and other uniquely microbial lipids, such as glycerol dialkyl
24 glycerol tetraether (GDGT) lipids, is thus of widespread interest but has so far not been

25 possible due to their size which prohibits analysis by traditional gas chromatography
26 pyrolysis isotope ratio mass spectrometry (GC-P-IRMS).

27 Methods: We determined the $\delta^2\text{H}$ values of large, polar molecules and applied high
28 temperature gas chromatography (GC) methods on a modified GC-P-IRMS system. The
29 methods were validated using authentic standards of large, functionalised molecules
30 (triacylglycerides, TAG), purified reference standards of GDGTs, and compared to $\delta^2\text{H}$
31 values determined by elemental analyser pyrolysis isotope ratio mass spectrometry (EA-P-
32 IRMS); and subsequently applied to the analysis of GDGTs in a sample from a methane
33 seep and a Welsh peat.

34 Results: $\delta^2\text{H}$ values of TAGs agreed within error between different between GC-P-IRMS and
35 EA-P-IRMS, with GC-P-IRMS showing 3-5 ‰ precision for 10 ng H injected. Archaeal lipid
36 GDGTs with up to three cyclisations could be analysed: $\delta^2\text{H}$ values were not significantly
37 different between methods with standard deviations of 5 to 6 ‰. When environmental
38 samples were analysed, $\delta^2\text{H}$ values of isoGDGTs were 50 ‰ more negative than those of
39 terrestrial brGDGTs.

40 Conclusions: Our results indicate that the high temperature GC-P-IRMS (HTGC-P-IRMS)
41 method developed here is appropriate to determine the $\delta^2\text{H}$ values of TAGs, GDGT lipids
42 with up to two cyclisations, and potentially other high molecular weight compounds. The
43 methodology will widen the current analytical window for biomarker and alimentary light
44 stable isotope analyses. Moreover, our initial measurements suggest that bacterial and
45 archaeal GDGT $\delta^2\text{H}$ values can record environmental and ecological conditions.

46

47 **Introduction**

48 The stable hydrogen isotopic composition ($\delta^2\text{H}$ values) of water varies systematically across
49 the globe ¹⁻³. The $\delta^2\text{H}$ values of biological molecules, in turn, are dependent on the $\delta^2\text{H}$ of
50 the H_2O available to the producing organism (source water), overprinted by biochemical
51 processes. $\delta^2\text{H}$ values of bulk organic matter and individual compounds are used across a
52 range of disciplines, e.g., in ecology and biology to trace animal migration patterns and
53 foodwebs ^{4,5}, in forensic science to identify geographical origins of victims or suspects ⁶, and
54 in food science to determine the provenance of products such as honey ⁷, milk ⁸, and meat ⁹.
55 The determination of $\delta^2\text{H}$ values has also resulted in substantial discoveries in archaeology,
56 such as the earliest horse milking ¹⁰, or manuring practices ¹¹, and has improved our
57 understanding of past environments and precipitation regimes ¹²⁻¹⁴.

58 The $\delta^2\text{H}$ values of individual lipid biomarkers are particularly useful in paleoenvironmental
59 studies. In particular, the correlation of lipid $\delta^2\text{H}$ with source water $\delta^2\text{H}$ has been widely
60 documented ^{12,15,16}, such that leaf waxes are now widely used to reconstruct past
61 hydrological conditions ^{12,16-18}. Long-chain *n*-alkanes and other non-functionalised
62 hydrocarbons are often used, because they are inherently less susceptible to hydrogen
63 exchange than other compound classes commonly found in sediments, due to their
64 prohibitively high pK_as (~ 50). However, a wide range of sedimentary lipids have been
65 analysed for their stable hydrogen isotopic composition, including *n*-alkanes, fatty acids,
66 alkenones, and, to a lesser extent, sterols and hopanols ¹⁹⁻²³.

67 The routine and rapid compound-specific $\delta^2\text{H}$ value determination of biomarkers (as
68 opposed to labour intensive approaches requiring compound isolation and purification)
69 requires the application of gas chromatography, coupled to an in-line reactor containing
70 active graphite, converting individual organic compounds into CO and H_2 ^{21,24-27}. The
71 produced gas is introduced into a mass spectrometer detecting *m/z* 2 (H-H) and 3 (H-D).
72 This setup requires analytes to be GC-amenable ²⁸, limiting analyses to compounds of a
73 molecular weight and polarity low enough to elute at a typical maximum capillary column

74 operating temperature of 320 °C. Therefore, only very few larger compounds (eluting later
75 than a C₃₆ *n*-alkane on an apolar stationary phase) have had their δ²H values successfully
76 determined. Existing measurements were achieved by implementing long isothermal holds
77 at 320 °C but only with highly purified and ²H-labelled compounds ²⁹, due to the low GC
78 resolution and δ²H precision associated with this methodology.

79 However, the δ²H values of large and/or polar compounds can be of significant interest. For
80 example, the origin of vegetable oils and milk products can be constrained ³⁰⁻³² with greater
81 specificity when isotopic fingerprinting is based on individual fatty acids instead of bulk
82 organics ^{33,34}. Moreover, determining the δ²H values of intact triacylglycerides (TAG, Suppl.
83 Fig. 1A), instead of hydrolysed and derivatised fatty acids, could have many benefits such as
84 eliminating derivatisation biases and increased specificity. TAG are routinely characterised in
85 food forensics by high temperature gas chromatography (HTGC; Buchgraber et al., 2004;
86 Fontecha et al., 2006; Ruiz-Samblas et al., 2015), but their ²H signatures are yet to be
87 exploited. Another potential application arises from very long-chain *n*-alkanes that are major
88 constituents of crude oil; their δ²H values could be used to assess source rock potential
89 ^{17,18,38,39}, or for correlating different oils and source rocks ^{38,40}.

90 A third suite of applications centres on glycerol dialkyl glycerol tetraether lipids (GDGTs,
91 Suppl. Fig. 1BC), derived from both Archaea and Bacteria and of wide interest in
92 geochemistry. These membrane lipids are frequently used in proxies for paleotemperature
93 and other environmental variables (reviewed in Schouten et al., 2013). In many sedimentary
94 archives, GDGTs are of mixed origins (e.g. De Jonge et al., 2014; Peterse et al., 2009), and
95 their δ²H values could thus be used to distinguish terrigenous from in situ-produced GDGTs,
96 for example in marine and lacustrine sediments. This would substantially improve the
97 application of these GDGT-based proxies. Moreover, in single-source environments, the
98 hydrogen isotopic composition of GDGTs could serve as a paleohydrological proxy, enabling
99 reconstruction of salinity, elevation, or precipitation. More recently, it has been shown that
100 δ²H values of bacterial lipids document the metabolic state of the source organisms,

101 potentially representing another application in biogeochemical investigations (Wijker et al.,
102 2019), and this method will allow to extend such investigations to Archaea.

103 In order to determine the stable isotopic composition of some of these large molecules, they
104 are often subjected to chemical degradation, and only fragments (mostly aliphatic moieties)
105 that are more GC-amenable than the parent molecule are analysed by GC-IRMS. For
106 TAGs, this involves acid methanolysis⁴⁴. For GDGTs, this involves ether cleavage, followed
107 by reduction⁴⁵⁻⁵⁰, often including laborious preparative HPLC steps for cleaning and
108 preconcentration⁵¹. Aside from being labour intensive, such procedures under acidic
109 conditions can result in hydrogen exchange.

110 However, recently, high temperature GC methods for more direct analysis of these
111 compounds have been developed; identification and quantification of GDGTs has been
112 achieved employing HTGC coupled to time-of-flight mass spectrometry (HTGC-TOFMS) and
113 flame ionisation detection (HTGC-FID^{52,53}). Here, we develop these methods further and
114 demonstrate $\delta^2\text{H}$ analysis of polar and high molecular weight compounds by high
115 temperature gas chromatography coupled to pyrolysis isotope ratio mass spectrometry
116 (HTGC-P-IRMS). We compare the values of purchased, authentic standards (TAGs), and
117 purified standards (GDGTs) determined by elemental analyser pyrolysis isotope ratio mass
118 spectrometry (EA-P-IRMS) with the values determined by HTGC-P-IRMS. We then report
119 the $\delta^2\text{H}$ values of GDGTs in a number of environmental samples.

120 **Experimental**

121 *Standards and environmental samples*

122 Triacylglyceride [trimyristin (C_{42}), tripalmitin (C_{48}), and tristearin (C_{54})] and *n*-alkane
123 standards were purchased from Sigma Aldrich (Gillingham, UK). isoGDGT-2 and isoGDGT-3
124 standards were purified from biomass of *Sulfolobus solfataricus* (DSM 1616), which was
125 grown in two batches (2 L each) of modified Allen medium⁵⁴ using water with a $\delta^2\text{H}$ value of
126 -55.0 ± 0.2 ‰. Each batch was inoculated with 20 ml of a late log-phase culture, incubated

127 aerobically at 76 °C with agitation at 200 RPM, and harvested in mid-log phase at an optical
128 density of 0.442 (600 nm). Cells were collected by centrifugation at 4 °C, frozen in liquid
129 nitrogen, and freeze-dried. 0.5 g of the freeze-dried cell pellet was subjected to acid
130 hydrolysis in 5 mL of 1.5 N methanolic HCl (10 % H₂O), and lipids were extracted by
131 ultrasonication in dichloromethane:methanol (1:1; v/v) as previously described⁵⁵. The total
132 lipid extract (TLE) was dried under a stream of N₂, dissolved in 1 mL of *n*-hexane:
133 isopropanol (97:3; v/v), and filtered through a 0.45 µm PTFE filter.

134 Individual isoprenoidal GDGTs containing 2 and 3 cyclopentyl moieties (isoGDGT-2 and
135 isoGDGT-3) were isolated by preparative normal phase (NP) high-performance liquid
136 chromatography (HPLC). To this end, aliquots (25 µL) of the filtered TLE were injected onto
137 an Agilent 1100 HPLC system fitted with an Econosphere NH₂ column (250 × 10 mm,
138 10 µm; Grace/Alltech). GDGTs were eluted isocratically with a solvent mixture of 1.35 %
139 isopropanol (IPA) in *n*-hexane at a flow rate of 1 mL min⁻¹ for 45 min, and the column was
140 cleaned with 16 % IPA for 12 min and re-equilibrated to initial conditions for 13 min after
141 every run. GDGTs were recovered by time-based fraction collection, according to the elution
142 times determined by atmospheric pressure chemical ionisation-mass spectrometry (APCI-
143 MS) using an Agilent 1100 MSD⁵⁶. The collected fractions were analysed by flow injection
144 analysis-mass spectrometry on the same instrument, and subsequently pooled by
145 compound. The purity of each isolated GDGT was >97 % as assessed by NP and reverse
146 phase HPLC-ACPI-MS analysis of the combined fractions⁵⁷, scanning the range *m/z* of
147 350–1350.

148 Environmental samples analysed included a sediment sample from a marine methane seep,
149 and a sample from a Welsh peat⁵². In order to improve gas chromatographic performance,
150 GDGTs were purified prior to HTGC-P-IRMS. The Welsh peat extract was passed over a
151 column containing 130-270 mesh silica (pore size 60 Å 288608, Sigma Aldrich, Gillingham,
152 UK) conditioned in methanol, using two column volumes of each hexane,
153 ethylacetate/hexane 1:9 (v/v), 25:75, 50:50, pure ethylacetate, and methanol.

154 Concentrations of GDGTs in the fractions were confirmed by adding triglyceride
155 quantification standards and analysis by HTGC-FID⁵². All fractions containing GDGTs
156 (Suppl. Fig. 2) were combined to avoid any isotope fractionation which may have occurred
157 during column chromatography.

158 ²H analysis by EA-P-IRMS

159 The ²H/¹H ratios of the triacylglycerides (TAGs) and C₅₀ and C₆₀ *n*-alkanes were analysed by
160 EA-P-IRMS at Elementar UK Ltd (EUK; Stockport, UK) and University of Colorado (CUB;
161 Boulder, USA). CUB also analysed GDGTs. CUB performed EA-P-IRMS analysis on a Flash
162 HT Plus elemental analyser with zero blank autosampler coupled to a Delta V Plus IRMS via
163 ConFlo-IV Interface (Thermo Scientific). At EUK, EA-P-IRMS measurements were performed
164 using a Geovision, which comprised a vario PYRO cube coupled to an isoprime vision IRMS.
165 Both laboratories measured samples using glassy carbon reactors in oxygen-free
166 environments, and performed multipoint calibrations using reference materials provided by
167 Arndt Schimmelmann (Indiana University, Bloomington, IN, USA) in order to standardise
168 determined δ²H values against the international reference Vienna Standard Mean Ocean
169 Water (VSMOW). CUB calibrated using 5α-androstane #3 (-293.2 ± 1.0 ‰), eicosanoic acid
170 methyl ester #Z1 / USGS 70 (-183.9 ± 1.4 ‰), and eicosanoic acid methyl ester #Z2 /
171 USGS 71 (-4.9 ± 1.0 ‰), and EUK calibrated using tetracosane #1: -53.0 per mil ± 1.6 ‰,
172 pentacosane #4: -263.6 ± 2.2 ‰ and heptacosane #3: -172.80 ± 1.6 ‰, and a standard
173 provided by the International Atomic Energy Agency, Vienna (IAEA CH-7: -100.2 ± 1.0 ‰).
174 Across both labs, the standard deviation (SD) of triplicate sample analyses was typically <
175 ±0.75 ‰.

176 Because the oxygen-bound H atoms of the GDGTs' hydroxyl moieties are easily exchanged,
177 the ²H content at these positions may have been altered during solvent
178 extraction/evaporation. We therefore vapour-equilibrated the dried GDGT fractions with local
179 deionised water (-121.8 ± 1.3 ‰) before analysis (24 h at 25 °C). GDGT fractions were then
180 dissolved in ethyl acetate at ~10 µg µL⁻¹ and 10 µL aliquots were pipetted into combusted

181 (450 °C, 10 h) silver capsules (4x6 mm), which were pre-loaded with small discs (d = 4 mm)
182 of combusted glass fibre filters (Whatman GF/F) as a solvent adsorbent. The solvent was
183 then completely evaporated in a closed chamber continuously purged with N₂ (30 min at
184 ~30 mL min⁻¹). Analysis by EA-P-IRMS was then conducted as described above.

185 To test for the efficiency of the vapour equilibration, a synthetic diglycerol-trialkyl-tetraether
186 (C₄₆-GTGT; Patwardhan and Thompson, 1999) was exposed to vapour of both ²H-enriched
187 water (7 atom % ²H) and local deionised water (24 h at 25 °C). Exposure to ²H-enriched
188 water vapour increased the ²H content of the molecule by 0.1 atom % (from 0.014 to
189 0.113 atom % relative to total H), corresponding to a ²H content of ~5 atom % at the OH
190 positions after exposure (assuming all exchange is localised to the hydroxyl moieties).
191 Exposure to natural water vapor, however, did not lead to a change in δ²H within analytical
192 precision of the measurement. The induced ²H content at the OH positions decreased again
193 to a ²H content of ~2 atom % at the OH-positions after a 12 h exposure to ambient lab air.
194 Together this indicates that OH-bound H of diglycerol tetraethers is readily exchanged with
195 ambient water vapor, and any ²H enrichment resulting from the evaporation of OH-
196 containing solvents (e.g. methanol) were likely diminished either by spontaneous re-
197 equilibration with ambient air, or by the latest through 24 h exposure to natural water vapor
198 in a desiccator as described above.

199 *δ²H value determination by high-temperature GC-P-IRMS*

200 Before analysis by HTGC-IRMS, fractions containing GDGTs and the sample from the Black
201 Sea methane seep were dissolved in 50 µl pyridine and derivatised to trimethylsilylethers
202 with 50 µl 99% N,O-Bis(trimethylsilyl)trifluoroacetamide (BSTFA), 1% trimethylchlorosilane
203 (TMCS), for one hour at 70 °C. The δ²H value of the TMS moieties used to derivatise the
204 hydroxyl-groups (δ²H_{TMS}) was determined by derivatisation of sodium palmitate δ²H_P of a
205 known δ²H, and analysis by GC-IRMS to yield the values of derivatised palmitate δ²H_{TMSP}, as
206 -82.35 ‰ acc. to Eqn. 1.

207
$$\delta^2 H_{TMS} = \frac{\delta^2 H_{TMSP} \cdot 40 - \delta^2 H_P}{9}$$

208 (Eqn 1)

209 Values of derivatised GDGTs $\delta^2 H_{meas}$ were corrected by mass balance to give $\delta^2 H_{GDGT}$ with n
210 representing the number of non-exchangeable hydrogens of the compounds and k the
211 number of TMS groups added (1 for archaeol, 2 for GDGTs and hydroxyarchaeol; Eqn. 2).

212
$$\delta^2 H_{GDGT} = \frac{\delta^2 H_{meas}(n + k \cdot 9)}{n} - \frac{k \cdot 9 \cdot \delta^2 H_{TMS}}{n}$$

213 (Eqn 2)

214 This was combined into Eqn. 3.

215
$$\delta^2 H_{GDGT} = \frac{\delta^2 H_{meas}(n + k \cdot 9)}{n} - \frac{k \cdot 40 \cdot \delta^2 H_{TMSP}}{n} + \frac{k \cdot 31 \cdot \delta^2 H_P}{n}$$

216 (Eqn 3)
217

218 Errors of $\delta^2 H_{meas}$ were determined according to error propagation laws:

219
$$\sigma_{\delta^2 H_{GDGT}}^2 = \sigma_{\delta^2 H_{meas}}^2 \cdot \left(\frac{n + k \cdot 9}{n}\right)^2 + \sigma_{\delta^2 H_{TMSP}}^2 \cdot \left(\frac{k \cdot 40}{n}\right)^2 + \sigma_{\delta^2 H_P}^2 \cdot \left(\frac{k \cdot 31}{n}\right)^2$$

220 (Eqn 4)

221 Samples were screened by HTGC-FID as described by Lengger et al.⁵² before they were
222 analysed by an Elementar isoprime visION HTGC-P-IRMS (Elementar UK Ltd., Cheadle,
223 UK). The instrument comprised an Agilent 7890B GC fitted with an on-column injector, linked
224 to a GC5 interface (maintained at 380 °C) and a hollow ceramic reactor, enabling pyrolysis
225 at 1450 °C. Ferrules used to connect the ceramic furnace and GC-column, as well as the
226 sample line He used as an additional carrier in the GC-IRMS system, were 100% graphite.
227 Ion beams at m/z 2 and 3 were monitored via an isoprime visION mass spectrometer. The
228 H_3^+ factor was determined daily or at least every 4 runs. Compounds were separated on a
229 Zebtron ZB-5HT analytical column (7 m × 0.25 mm × 0.1 µm) with high-temperature resistant
230 polyimide coating, which was fitted to a transfer line and an exhaust to allow diversion of the
231 solvent peak to waste via a glass Y-splitter fixed with high temperature resin (Phenomenex
232 Ltd., Aschaffenburg, Germany). He was used as a carrier gas at a flow rate of 2.2 ml min⁻¹,

233 and the oven was programmed as follows: 1 min hold at 70 °C, increase by 10 °C min⁻¹ to
234 350 °C, followed by an increase at 3 °C min⁻¹ to 400 °C (10 min hold). Results were
235 calibrated using a mixture of *n*-alkanes (B3, A. Schimmelmann, Indiana University,
236 Bloomington, IN, USA) according to Sessions et al.^{21,59}, which was injected at least every
237 four analyses, and analysed using a He flow of 1 ml min⁻¹, with a different temperature
238 program (injection at 50 °C held for 1 min followed by an increase of 10 °C min⁻¹ to 300 °C
239 and a 10 min hold). Resultant calibrated δ²H values were calculated based on the derived
240 linear regression. Root mean standard errors of normalised values of the *n*-alkanes mixture
241 were typically between 4 and 6 ‰, and never exceeded 10 ‰. Data was processed using
242 ionOS stable isotope data processing software (Elementar UK Ltd., UK), using an automated
243 multi-point linearisation based on the certified values of the 15 individual *n*-alkanes
244 comprising the B3 standard.

245 The fractionation factor ε_{H₂O/GDGT} was determined from the δ²H_{H₂O} and the δ²H_{GDGT} (Eqn. 5).

$$246 \quad \varepsilon_{H_2O/GDGT} = \left(\frac{\delta_{H_2O} + 1}{\delta_{GDGT} + 1} - 1 \right) \cdot 1000$$

247 (Eqn. 5)

248 **Results and discussion**

249 *Chromatographic resolution*

250 The modifications of the GC-IRMS setup enabled operating temperatures of up to 400 °C.
251 Utilisation of a 7-m column and on-column injection enabled elution of isoGDGTs up to
252 GDGT-3, as well as acceptable values for the B3 standard. The GC-IRMS required a
253 polyimide-coated column rather than the metal column commonly employed in high
254 temperature GC-methodologies, as this allowed flow diversion via a glass Y-splitter in which
255 the column was secured using high temperature resin. The glass Y-splitter ensured minimal
256 thermal mass. Furthermore, the pneumatically operated heart-cut valve enabling diversion of
257 the solvent away from the furnace reactor was moved to a location outside of the GC-oven in

258 order to avoid potential leaks associated with the high temperatures. Methods employing
259 15 and 30 m polyimide-coated capillary columns, equivalent to the metal columns that had
260 been successfully used to analyse isoGDGTs by HTGC-FID and HTGC-TOFMS at extended
261 > 400 °C isothermals, could not be employed to elute isoGDGTs in analogous HTGC-P-
262 IRMS analyses due to their comparatively low stability at these temperatures.

263 The unusual HTGC configuration, with a short 7 m column, high flow, and on-column
264 injector, was tested by analysing a mixture of 15 *n*-alkanes: the so-called Indiana B-standard
265 mix routinely used for standardisation of GC-IRMS results. Baseline separation of individual
266 *n*-alkane peaks and acceptable root mean square errors were achieved with this method
267 (Fig. 1A): this standard was subsequently used for quality control and isotope calibration.
268 Root mean square error (RMSE) and linearisation equations for all analyses of the standards
269 are given in Supplementary Fig. 3 and Table 1, with linearisation applied to the samples
270 based on the most contemporary analysis of the standard. RMSE for all accepted analyses
271 were always below 10 ‰: whenever 10 was exceeded, inlet maintenance or column
272 changes were performed. An *n*-alkane standard containing higher molecular weight
273 compounds (up to C₆₀, Fig. 1B), a mixture of triacylglycerides (Fig. 1C), a seep sample
274 containing GDGT-0, -1, -2, and -3, and the two GDGT standards (GDGT-2 and -3) (Fig. 1D)
275 were analysed and chromatograms were similar to previous results employing HTGC-FID
276 and a 7 m column⁵². The brGDGTs eluted earlier than isoGDGTs (cf. ⁵²).

277 *Accuracy and precision of δ²H values of high molecular weight compounds*

278 Purchased triacylglyceride (TAG) reference compounds and purified GDGT standards were
279 used to test the methodology for accuracy by determining the δ²H values of these
280 compounds by HTGC-IRMS at GC temperatures of up to 400 °C as well as by EA-analysis.
281 The prepared isoGDGT-2 and isoGDGT-3 standards were analysed by one laboratory (CU
282 Boulder), while the purchased standards were examined by EA-P-IRMS in two different
283 laboratories (CU Boulder and Elementar UK Ltd). The average δ²H values determined for
284 the TAGs were within 5 ‰ for all analyses (Tab. 1, Fig. 2). HTGC-analysed samples

285 generally yielded $\delta^2\text{H}$ values between the values determined by the EA analyses. Standard
286 deviations were smaller for the EA methods ($< 2 \text{ ‰}$) than for the HTGC method (9-18 ‰,
287 which represents 2-3 \times the typical precision of $\delta^2\text{H}$ value determinations by GC-IRMS;
288 Sessions, 2006). However, here injection concentrations varied, which likely contributed to
289 the variability, which we investigate further below. It is expected that further application of
290 this technique – and routine analysis of TAGs, as compounds of particular interest to the
291 alimentary industry – will lead to improvements in analytical precision as methods are
292 improved by optimising solvents, injection temperatures, and concentrations. The $\delta^2\text{H}$ values
293 determined for the high molecular weight *n*-alkanes with 50 and 60 carbon atoms (Table 1)
294 were more variable among all methods and laboratories. This was surprising, and possibly a
295 result of insufficient mixing of these large waxy compounds before distribution.

296 The $\delta^2\text{H}$ values of purified GDGTs obtained by EA-P-IRMS and HTGC-IRMS (Tab. 1) were
297 not significantly different for GDGT-2 at a high confidence level (Welsh's t-test, $df = 2$, $t =$
298 1.32 , $p = 0.32$). However, for GDGT-3, which eluted later, the $\delta^2\text{H}$ value derived by HTGC-
299 IRMS was 9 ‰ higher than the value determined by EA-P-IRMS ($df = 2$, $t = 3.32$, $p = 0.080$).
300 A raised baseline could be a possible cause for this discrepancy, but ionOS software applies
301 an automated correction, and both GDGTs eluted on an isothermal baseline. A more likely
302 cause could be fractionation due to chromatographic separation, to adsorption cold spots, or
303 thermal decomposition. Another possibility is minor contamination of GDGT-3, resulting in a
304 flawed EA-P-IRMS measurement but not affecting HTGC-P-IRMS measurements; however,
305 this would be surprising as GDGT-2 and GDGT-3 were isolated from the same organism and
306 the EA-P-IRMS results match expectations of similar $\delta^2\text{H}$ values. The standard deviation of 5
307 – 6 ‰ achieved for purified GDGTs using the HTGC-P-IRMS system is similar to the
308 precision of lower molecular weight compounds on a conventional GC-P-IRMS instrument
309 (e.g. Sessions, 2006).

310 *Response vs accuracy*

311 Whilst GDGTs are ubiquitous, they are typically only present at ppm to ppb concentrations in
312 environmental samples such as sediments and soils. In addition, many high molecular
313 weight compounds are not very soluble in solvents suitable for GC-IRMS, and on-column
314 injection only allows small amounts of sample to be used. Therefore, only small amounts of
315 GDGT (ng) were injected for each HTGC-IRMS analysis. To assess accuracy in relationship
316 to signal intensity, different concentrations of the TAG standard were tested and compared
317 to peak heights (Fig. 3). This yielded a response of 0.07 – 0.08 nA per ng H per compound
318 for m/z 2. Below ~0.25 nA peak height, values begin to deviate substantially (by ~20 ‰) from
319 the values measured by EA-P-IRMS, with differences of up to 400 ‰ when peak heights
320 were around 0.1 nA. We thus excluded peak heights < 0.25 nA, corresponding to less than
321 3.5 ng H injected on column. Typical H amounts required to achieve 3-5 ‰ precision were
322 ~10 ng, translating to m/z 2 peak heights of 0.7 – 0.8 nA (equivalent to 700-800 mV on an
323 IRMS with a 10^9 Ohm resistor on the operational amplifier for the m/z 2 faraday cup).

324 *GDGTs in environmental samples and $\epsilon_{H2O/isoGDGT}$*

325 A sample from a Mediterranean cold seep⁵² was analysed, and δ^2H values for archaeol,
326 hydroxyarchaeol, GDGT-1, and GDGT-2 were determined to be -245 ± 7 , -253 ± 13 , $-216 \pm$
327 15 , and -225 ± 14 , respectively ($n=3$; Fig. 1D, Fig. 4). These values show a limited range, as
328 expected for ether lipids derived from a common archaeal source, and are similar to
329 published δ^2H values of the biphytanes of GDGTs in *Sulfolobus sp.* determined after ether
330 cleavage (-229 to -257 ‰⁴⁵). However, the values are not identical, with the diphytanyl
331 glycerol diether lipids archaeol and hydroxyarchaeol being 2H -depleted relative to GDGTs.
332 Though the difference is small, it could potentially reflect different archaeal origins, given that
333 ANME-2 group Archaea appear to preferentially produce GDGTs in cold seep settings (e.g.,
334 Blumenberg et al., 2004); this would be particularly true if the differing source Archaea
335 exhibit different metabolisms (see below).

336 The $\epsilon_{H2O/GDGT}$ for the *Sulfolobus* cultures used to purify the standards was determined
337 as -134 ‰ and was lower than previously reported $\epsilon_{H2O/GDGT}$ (-213 ‰ to -161 ‰⁴⁵). The

338 application of this fractionation factor to the environmental iso-GDGTs would result in an
339 unrealistic $\delta^2\text{H}$ value for the seawater of -93‰ , suggesting that metabolism, salinity,
340 temperature, and other factors contribute strongly to the extent of fractionation.

341 Values for GDGT-0 from the peat (Suppl. Fig. 4) were similar to the isoGDGTs in the seep
342 sample ($-235 \pm 3\text{‰}$, $n = 2$), whereas values for brGDGTs were relatively enriched in ^2H
343 ($-176 \pm 6\text{‰}$, $n = 6$). It is possible that the ^2H -enrichment of brGDGTs relative to co-occurring
344 isoGDGTs could be due to fractionation associated with the biosynthetic pathways for
345 isoprenoidal (isoGDGTs) vs. *n*-acyl lipids (brGDGTs), in which isoprenoidal lipids (which
346 undergo successive hydrogenation) exhibit more ^2H -depleted signatures^{21,62}. However,
347 recently, it has also been shown that the energy and metabolism pathways of source
348 organisms are highly correlated with $\delta^2\text{H}$ values of their lipids^{63–65}; it is also thought that
349 NADPH/NADH ratios and transhydrogenases play an important role, particularly in
350 anaerobic organisms^{66–69}. In general, heterotrophic bacteria consuming TCA-cycle
351 intermediates exhibit $\delta^2\text{H}$ values similar to or enriched relative to source water, heterotrophs
352 assimilating carbohydrates are slightly depleted relative to source water, and
353 photoautotrophic and chemoautotrophic bacteria show the greatest ^2H -depletion⁶⁴. While
354 Archaea were not examined in this work, some of our results are consistent with the idea
355 that chemoautotrophic archaea are the presumed producers of isoGDGTs in both settings,
356 and heterotrophic bacteria are thought to be the producers of brGDGTs⁷⁰.

357 The differences between the peat and seep samples for isoGDGTs are unexpected: As the
358 $\delta^2\text{H}$ of the peat water is likely around -52‰ ¹ – slightly depleted compared to seawater – we
359 expected isoGDGTs from peat to also be depleted in ^2H relative to GDGTs from marine
360 environments. However, isoGDGTs from peat are up to 10 to 20 ‰ more ^2H -enriched in the
361 peat, invoking a difference in metabolic state between the anaerobic methanogens in peat,
362 and the anaerobic methane oxidising communities in the seep. These findings speak to the
363 potential of isoGDGT $\delta^2\text{H}$ analyses in probing microbial ecology and metabolic state, while

364 brGDGTs, which are presumably of heterotrophic bacterial origin in peat settings, could
365 prove useful as proxies for source water $\delta^2\text{H}$ and hydrology.

366 The novel HTGC-P-IRMS method enables the determination of the $\delta^2\text{H}$ values of
367 compounds with a high molecular weight, including TAG and GDGT lipids, hereby extending
368 the range of analytes for $\delta^2\text{H}$ value determination. Accuracy and precision are as small as
369 3 ‰ in some cases and comparable to EA-P-IRMS. Our initial measurements suggest that
370 bacterial and archaeal GDGT $\delta^2\text{H}$ values are likely related to both environmental
371 parameters, and the metabolic and ecological function of the source organisms. Future
372 applications include but are not limited to alimentary forensics, archaeology, oil-source rock
373 correlations, microbial ecology and paleoclimate.

374 **Acknowledgements**

375 The authors would like to thank Paul Sutton, Alison Kuhl, Hanna Gruszczynska, Ed Aldred,
376 Xiahong Feng, Alec Cobban, and Michiel Kienhuis for support with measurements, advice,
377 and discussions of techniques. SKL was funded by a Rubicon Grant 825.14.014 from the
378 Netherlands Organisation for Scientific Research (NWO). RDP acknowledges support from
379 ERC (Advanced Grant T-GRES, to RDP). AP and YW acknowledge support from the Swiss
380 National Science Foundation (P2BSP2_168716), and from the Gordon and Betty Moore
381 Foundation and US National Science Foundation (to AP). SHK acknowledges support from
382 the US National Science Foundation. The authors thank the Natural Environment Research
383 Council, UK, for partial funding of the mass spectrometry facilities at Bristol (contract no.
384 R8/H10/63). WDL acknowledges support from the American Chemical Society (PRF 57209-
385 DNI2).

386

Table 1. $\delta^2\text{H}$ values determined by EA-IRMS and GC-IRMS.

	EA-IRMS (Elementar)			EA-IRMS (CU Boulder)			HTGC-IRMS		
	Mean	St. dev.	N	Mean	St. dev.	N	Mean	St. dev.	N
	[‰ V-SMOW]			[‰ V-SMOW]			[‰ V-SMOW]		
GDGT-2	-	-	-	-181.61	0.41	3	-186.5	6.4	3
GDGT-3	-	-	-	-182.6	0.19	3	-173.0	5.0	3
C₄₂-TAG	-234.97	0.48	4	-237.97	0.72	3	-232.8	9.2	9
C₄₈-TAG	-219.43	1.76	3	-224.14	0.33	3	-222.9	18.2	7
C₅₄-TAG	-225.85	0.38	4	-228.16	0.7	3	-223.0	11.6	7
<i>n</i>-C₆₀ alk	-206.49	4.65	3	-213.98	0.35	3	-196.3	2.6	3
<i>n</i>-C₅₀ alk	-199.29	0.02	4	-202.05	0.06	3	-187.5	3.0	3

References

1. Bowen GJ, Revenaugh J. Interpolating the isotopic composition of modern meteoric precipitation. *Water Resour Res.* 2003;39(10). doi:10.1029/2003WR002086
2. Craig H. Isotopic Variations in Meteoric Waters. *Science.* 1961;133(3465):1702-1703. doi:10.1126/science.133.3465.1702
3. West JB, Bowen GJ, Dawson TE, Tu KP, eds. *Isoscapes*. Springer Netherlands; 2010. doi:10.1007/978-90-481-3354-3
4. Hobson KA, Wassenaar LI. *Tracking Animal Migration with Stable Isotopes*. Academic Press; 2008.
5. Soto DX, Wassenaar LI, Hobson KA. Stable hydrogen and oxygen isotopes in aquatic food webs are tracers of diet and provenance. *Funct Ecol.* 2013;27(2):535-543. doi:10.1111/1365-2435.12054
6. Fraser I, Meier-Augenstein W. Stable ^2H isotope analysis of modern-day human hair and nails can aid forensic human identification. *Rapid Commun Mass Spectrom.* 2007;21(20):3279-3285. doi:10.1002/rcm.3209
7. Schellenberg A, Chmielus S, Schlicht C, et al. Multielement stable isotope ratios (H, C, N, S) of honey from different European regions. *Food Chem.* 2010;121(3):770-777. doi:10.1016/j.foodchem.2009.12.082
8. Chesson LA, Valenzuela LO, O'Grady SP, Cerling TE, Ehleringer JR. Hydrogen and Oxygen Stable Isotope Ratios of Milk in the United States. *J Agric Food Chem.* 2010;58(4):2358-2363. doi:10.1021/jf904151c
9. Chesson LA, Podlesak DW, Erkkila BR, Cerling TE, Ehleringer JR. Isotopic consequences of consumer food choice: Hydrogen and oxygen stable isotope ratios in foods from fast food restaurants versus supermarkets. *Food Chem.* 2010;119(3):1250-1256. doi:10.1016/j.foodchem.2009.07.046
10. Outram AK, Stear NA, Bendrey R, et al. The Earliest Horse Harnessing and Milking. *Science.* 2009;323(5919):1332-1335. doi:10.1126/science.1168594
11. Reynard LM, Hedges REM. Stable hydrogen isotopes of bone collagen in palaeodietary and palaeoenvironmental reconstruction. *J Archaeol Sci.* 2008;35(7):1934-1942. doi:10.1016/j.jas.2007.12.004
12. Sachse D, Billault I, Bowen GJ, et al. Molecular Paleohydrology: Interpreting the Hydrogen-Isotopic Composition of Lipid Biomarkers from Photosynthesizing Organisms. *Annu Rev Earth Planet Sci.* 2012;40(1):221-249. doi:10.1146/annurev-earth-042711-105535
13. Schefuß E, Schouten S, Schneider RR. Climatic controls on central African hydrology during the past 20,000 years. *Nature.* 2005;437(7061):1003-1006. doi:10.1038/nature03945
14. Tierney JE, Russell JM, Huang Y, Damsté JSS, Hopmans EC, Cohen AS. Northern Hemisphere Controls on Tropical Southeast African Climate During the Past 60,000 Years. *Science.* 2008;322(5899):252-255. doi:10.1126/science.1160485

15. Estep MF, Hoering TC. Biogeochemistry of the stable hydrogen isotopes. *Geochim Cosmochim Acta*. 1980;44(8):1197-1206. doi:10.1016/0016-7037(80)90073-3
16. da Silveira Lobo Sternberg L. D/H ratios of environmental water recorded by D/H ratios of plant lipids. *Nature*. 1988;333(6168):59-61. doi:10.1038/333059a0
17. Sessions AL. Review: Factors Controlling the Deuterium Contents of Sedimentary Hydrocarbons. *Org Geochem*. 2016;96:43-64. doi:10.1016/j.orggeochem.2016.02.012
18. Sessions AL, Sylva SP, Summons RE, Hayes JM. Isotopic exchange of carbon-bound hydrogen over geologic timescales. *Geochim Cosmochim Acta*. 2004;68(7):1545-1559. doi:10.1016/j.gca.2003.06.004
19. Li C, Sessions AL, Kinnaman FS, Valentine DL. Hydrogen-isotopic variability in lipids from Santa Barbara Basin sediments. *Geochim Cosmochim Acta*. 2009;73(16):4803-4823. doi:10.1016/j.gca.2009.05.056
20. Sauer PE, Eglinton TI, Hayes JM, Schimmelmann A, Sessions AL. Compound-specific D/H ratios of lipid biomarkers from sediments as a proxy for environmental and climatic conditions. *Geochim Cosmochim Acta*. 2001;65(2):213-222. doi:10.1016/S0016-7037(00)00520-2
21. Sessions AL, Burgoyne TW, Schimmelmann A, Hayes JM. Fractionation of hydrogen isotopes in lipid biosynthesis. *Org Geochem*. 1999;30(9):1193-1200. doi:10.1016/S0146-6380(99)00094-7
22. Smittenberg RH, Sachs JP. Purification of dinosterol for hydrogen isotopic analysis using high-performance liquid chromatography–mass spectrometry. *J Chromatogr A*. 2007;1169(1-2):70-76. doi:10.1016/j.chroma.2007.09.018
23. van der Meer MTJ, Baas M, Rijpstra WIC, et al. Hydrogen isotopic compositions of long-chain alkenones record freshwater flooding of the Eastern Mediterranean at the onset of sapropel deposition. *Earth Planet Sci Lett*. 2007;262(3-4):594-600. doi:10.1016/j.epsl.2007.08.014
24. Burgoyne TW, Hayes JM. Quantitative Production of H₂ by Pyrolysis of Gas Chromatographic Effluents. *Anal Chem*. 1998;70(24):5136-5141. doi:10.1021/ac980248v
25. Hilkert AW, Douthitt CB, Schlüter HJ, Brand WA. Isotope ratio monitoring gas chromatography/Mass spectrometry of D/H by high temperature conversion isotope ratio mass spectrometry. *Rapid Commun Mass Spectrom*. 1999;13(13):1226-1230. doi:10.1002/(SICI)1097-0231(19990715)13:13<1226::AID-RCM575>3.0.CO;2-9
26. Scrimgeour CM, Begley IS, Thomason ML. Measurement of deuterium incorporation into fatty acids by gas chromatography/isotope ratio mass spectrometry. *Rapid Commun Mass Spectrom*. 1999;13(4):271-274. doi:10.1002/(SICI)1097-0231(19990228)13:4<271::AID-RCM468>3.0.CO;2-6
27. Tobias HJ, Brenna JT. On-Line Pyrolysis as a Limitless Reduction Source for High-Precision Isotopic Analysis of Organic-Derived Hydrogen. *Anal Chem*. 1997;69(16):3148-3152. doi:10.1021/ac970332v

28. Kaal E, Janssen H-G. Extending the molecular application range of gas chromatography. *J Chromatogr A*. 2008;1184(1-2):43-60. doi:10.1016/j.chroma.2007.11.114
29. Bodlenner A, Liu W, Hirsch G, et al. C₃₅ Hopanoid Side Chain Biosynthesis: Reduction of Ribosylhopane into Bacteriohopanetetrol by a Cell-Free System from *Methylobacterium organophilum*. *ChemBioChem*. 2015;16(12):1764-1770. doi:10.1002/cbic.201500021
30. Muhammad SA, Seow E-K, Omar AM, et al. Variation of delta H-2, delta O-18 & delta C-13 in crude palm oil from different regions in Malaysia: Potential of stable isotope signatures as a key traceability parameter. *Sci Justice*. 2018;58(1):59-66. doi:10.1016/j.scijus.2017.05.008
31. Richter EK, Spangenberg JE, Kreuzer M, Leiber F. Characterization of Rapeseed (*Brassica napus*) Oils by Bulk C, O, H, and Fatty Acid C Stable Isotope Analyses. *J Agric Food Chem*. 2010;58(13):8048-8055. doi:10.1021/jf101128f
32. Spangenberg JE. Bulk C, H, O, and fatty acid C stable isotope analyses for purity assessment of vegetable oils from the southern and northern hemispheres. *Rapid Commun Mass Spectrom*. 2016;30(23):2447-2461. doi:10.1002/rcm.7734
33. Ehtesham E, Hayman AR, McComb KA, Van Hale R, Frew RD. Correlation of Geographical Location with Stable Isotope Values of Hydrogen and Carbon of Fatty Acids from New Zealand Milk and Bulk Milk Powder. *J Agric Food Chem*. 2013;61(37):8914-8923. doi:10.1021/jf4024883
34. Paolini M, Bontempo L, Camin F. Compound-specific delta C-13 and delta H-2 analysis of olive oil fatty acids. *Talanta*. 2017;174:38-43. doi:10.1016/j.talanta.2017.05.080
35. Buchgraber M, Ulberth F, Anklam E. Cluster analysis for the systematic grouping of genuine cocoa butter and cocoa butter equivalent samples based on triglyceride patterns. *J Agric Food Chem*. 2004;52(12):3855-3860. doi:10.1021/jf035153v
36. Fontecha J, Mayo I, Toledano G, Juárez M. Use of changes in triacylglycerols during ripening of cheeses with high lipolysis levels for detection of milk fat authenticity. *Int Dairy J*. 2006;16(12):1498-1504. doi:10.1016/j.idairyj.2005.12.005
37. Ruiz-Samblas C, Gonzalez-Casado A, Cuadros-Rodriguez L. Triacylglycerols Determination by High-temperature Gas Chromatography in the Analysis of Vegetable Oils and Foods: A Review of the Past 10 Years. *Crit Rev Food Sci Nutr*. 2015;55(11):1618-1631. doi:10.1080/10408398.2012.713045
38. Pedentchouk N, Turich C. Carbon and hydrogen isotopic compositions of n-alkanes as a tool in petroleum exploration. *Geol Soc Lond Spec Publ*. 2018;468(1):105-125. doi:10.1144/SP468.1
39. Radke J, Bechtel A, Gaupp R, et al. Correlation between hydrogen isotope ratios of lipid biomarkers and sediment maturity. *Geochim Cosmochim Acta*. 2005;69(23):5517-5530. doi:10.1016/j.gca.2005.07.014
40. Li M, Huang Y, Obermajer M, Jiang C, Snowdon LR, Fowler MG. Hydrogen isotopic compositions of individual alkanes as a new approach to petroleum correlation: case studies from the Western Canada Sedimentary Basin. *Org Geochem*. 2001;32(12):1387-1399. doi:10.1016/S0146-6380(01)00116-4

41. Schouten S, Hopmans EC, Sinninghe Damsté JS. The organic geochemistry of glycerol dialkyl glycerol tetraether lipids: A review. *Org Geochem*. 2013;54:19-61. doi:10.1016/j.orggeochem.2012.09.006
42. De Jonge C, Stadnitskaia A, Hopmans EC, Cherkashov G, Fedotov A, Sinninghe Damsté JS. In situ produced branched glycerol dialkyl glycerol tetraethers in suspended particulate matter from the Yenisei River, Eastern Siberia. *Geochim Cosmochim Acta*. 2014;125:476-491. doi:10.1016/j.gca.2013.10.031
43. Peterse F, Kim J-H, Schouten S, Kristensen DK, Koç N, Sinninghe Damsté JS. Constraints on the application of the MBT/CBT palaeothermometer at high latitude environments (Svalbard, Norway). *Org Geochem*. 2009;40(6):692-699. doi:10.1016/j.orggeochem.2009.03.004
44. Correa-Ascencio M, Evershed RP. High throughput screening of organic residues in archaeological potsherds using direct acidified methanol extraction. *Anal Methods*. 2014;6(5):1330-1340. doi:10.1039/C3AY41678J
45. Kaneko M, Kitajima F, Naraoka H. Stable hydrogen isotope measurement of archaeal ether-bound hydrocarbons. *Org Geochem*. 2011;42(2):166-172. doi:10.1016/j.orggeochem.2010.11.002
46. Lengger SK, Lipsewers YA, de Haas H, Sinninghe Damsté JS, Schouten S. Lack of ¹³C-label incorporation suggests low turnover rates of thaumarchaeal intact polar tetraether lipids in sediments from the Iceland shelf. *Biogeosciences*. 2014;11(2):201-216. doi:10.5194/bg-11-201-2014
47. Schouten S, Hoefs MJL, Koopmans MP, Bosch H-J, Sinninghe Damsté JS. Structural characterization, occurrence and fate of archaeal ether-bound acyclic and cyclic biphytanes and corresponding diols in sediments. *Org Geochem*. 1998;29(5-7):1305-1319. doi:10.1016/S0146-6380(98)00131-4
48. Weber Y, Damsté JSS, Zopfi J, et al. Redox-dependent niche differentiation provides evidence for multiple bacterial sources of glycerol tetraether lipids in lakes. *Proc Natl Acad Sci*. 2018;115(43):10926-10931. doi:10.1073/pnas.1805186115
49. Weber Y, De Jonge C, Rijpstra WIC, et al. Identification and carbon isotope composition of a novel branched GDGT isomer in lake sediments: Evidence for lacustrine branched GDGT production. *Geochim Cosmochim Acta*. 2015;154:118-129. doi:10.1016/j.gca.2015.01.032
50. Wegener G, Bausch M, Holler T, et al. Assessing sub-seafloor microbial activity by combined stable isotope probing with deuterated water and ¹³C-bicarbonate: Microbial carbon assimilation in the sub-seafloor. *Environ Microbiol*. 2012;14(6):1517-1527. doi:10.1111/j.1462-2920.2012.02739.x
51. Weijers JWH, Wiesenberg GLB, Bol R, Hopmans EC, Pancost RD. Carbon isotopic composition of branched tetraether membrane lipids in soils suggest a rapid turnover and a heterotrophic life style of their source organism(s). *Biogeosciences*. 2010;7(9):2959-2973. doi:10.5194/bg-7-2959-2010
52. Lengger SK, Sutton PA, Rowland SJ, et al. Archaeal and bacterial glycerol dialkyl glycerol tetraether (GDGT) lipids in environmental samples by high temperature-gas chromatography with flame ionisation and time-of-flight mass spectrometry detection. *Org Geochem*. 2018;121:10-21. doi:10.1016/j.orggeochem.2018.03.012

53. Sutton PA, Rowland SJ. High temperature gas chromatography–time-of-flight-mass spectrometry (HTGC–ToF-MS) for high-boiling compounds. *J Chromatogr A*. 2012;1243:69-80. doi:10.1016/j.chroma.2012.04.044
54. Worthington P, Blum P, Perez-Pomares F, Elthon T. Large-Scale Cultivation of Acidophilic Hyperthermophiles for Recovery of Secreted Proteins. *Appl Env Microbiol*. 2003;69(1):252-257. doi:10.1128/AEM.69.1.252-257.2003
55. Weber Y, Damsté JSS, Hopmans EC, Lehmann MF, Niemann H. Incomplete recovery of intact polar glycerol dialkyl glycerol tetraethers from lacustrine suspended biomass. *Limnol Oceanogr Methods*. 2017;15(9):782-793. doi:10.1002/lom3.10198
56. Schouten S, Hugué C, Hopmans EC, Kienhuis MVM, Sinninghe Damsté JS. Analytical methodology for TEX₈₆ paleothermometry by High-Performance Liquid Chromatography/Atmospheric Pressure Chemical Ionization-Mass Spectrometry. *Anal Chem*. 2007;79(7):2940-2944. doi:10.1021/ac062339v
57. Smittenberg RH, Hopmans EC, Schouten S, Sinninghe Damsté JS. Rapid isolation of biomarkers for compound specific radiocarbon dating using high-performance liquid chromatography and flow injection analysis–atmospheric pressure chemical ionisation mass spectrometry. *J Chromatogr A*. 2002;978(1):129-140. doi:10.1016/S0021-9673(02)01427-9
58. Patwardhan AP, Thompson DH. Efficient Synthesis of 40- and 48-Membered Tetraether Macrocyclic Bisphosphocholines. *Org Lett*. 1999;1(2):241-244. doi:10.1021/ol990567o
59. Sessions AL, Jahnke LL, Schimmelmann A, Hayes JM. Hydrogen isotope fractionation in lipids of the methane-oxidizing bacterium *Methylococcus capsulatus*. *Geochim Cosmochim Acta*. 2002;66(22):3955-3969. doi:10.1016/S0016-7037(02)00981-X
60. Sessions AL. Isotope-ratio detection for gas chromatography. *J Sep Sci*. 2006;29(12):1946-1961. doi:10.1002/jssc.200600002
61. Blumenberg M, Seifert R, Reitner J, Pape T, Michaelis W. Membrane lipid patterns typify distinct anaerobic methanotrophic consortia. *Proc Natl Acad Sci*. 2004;101(30):11111-11116. doi:10.1073/pnas.0401188101
62. Zhang Z, Sachs JP. Hydrogen isotope fractionation in freshwater algae: I. Variations among lipids and species. *Org Geochem*. 2007;38(4):582-608. doi:10.1016/j.orggeochem.2006.12.004
63. Osburn MR, Dawson KS, Fogel ML, Sessions AL. Fractionation of Hydrogen Isotopes by Sulfate- and Nitrate-Reducing Bacteria. *Front Microbiol*. 2016;7. doi:10.3389/fmicb.2016.01166
64. Wijker RS, Sessions AL, Fuhrer T, Phan M. ²H/¹H variation in microbial lipids is controlled by NADPH metabolism. *Proc Natl Acad Sci*. 2019;116(25):12173-12182. doi:10.1073/pnas.1818372116
65. Zhang X, Gillespie AL, Sessions AL. Large D/H variations in bacterial lipids reflect central metabolic pathways. *Proc Natl Acad Sci*. 2009;106(31):12580-12586. doi:10.1073/pnas.0903030106

66. Campbell BJ, Sessions AL, Fox DN, et al. Minimal Influence of [NiFe] Hydrogenase on Hydrogen Isotope Fractionation in H₂-Oxidizing Cupriavidus necator. *Front Microbiol.* 2017;8. doi:10.3389/fmicb.2017.01886
67. Dawson KS, Osburn MR, Sessions AL, Orphan VJ. Metabolic associations with archaea drive shifts in hydrogen isotope fractionation in sulfate-reducing bacterial lipids in cocultures and methane seeps. *Geobiology.* 2015;13(5):462-477. doi:10.1111/gbi.12140
68. Leavitt WD, Murphy SJ-L, Lynd LR, Bradley AS. Hydrogen isotope composition of Thermoanaerobacterium saccharolyticum lipids: Comparing wild type with a nfn-transhydrogenase mutant. *Org Geochem.* 2017;113:239-241. doi:10.1016/j.orggeochem.2017.06.020
69. Leavitt WD, Flynn TM, Suess MK, Bradley AS. Transhydrogenase and Growth Substrate Influence Lipid Hydrogen Isotope Ratios in Desulfovibrio alaskensis G20. *Front Microbiol.* 2016;07. doi:10.3389/fmicb.2016.00918
70. Sinninghe Damsté JS, Rijpstra WIC, Foesel BU, et al. An overview of the occurrence of ether- and ester-linked iso-diabolic acid membrane lipids in microbial cultures of the Acidobacteria: Implications for brGDGT paleoproxies for temperature and pH. *Org Geochem.* 2018;124:63-76. doi:10.1016/j.orggeochem.2018.07.006

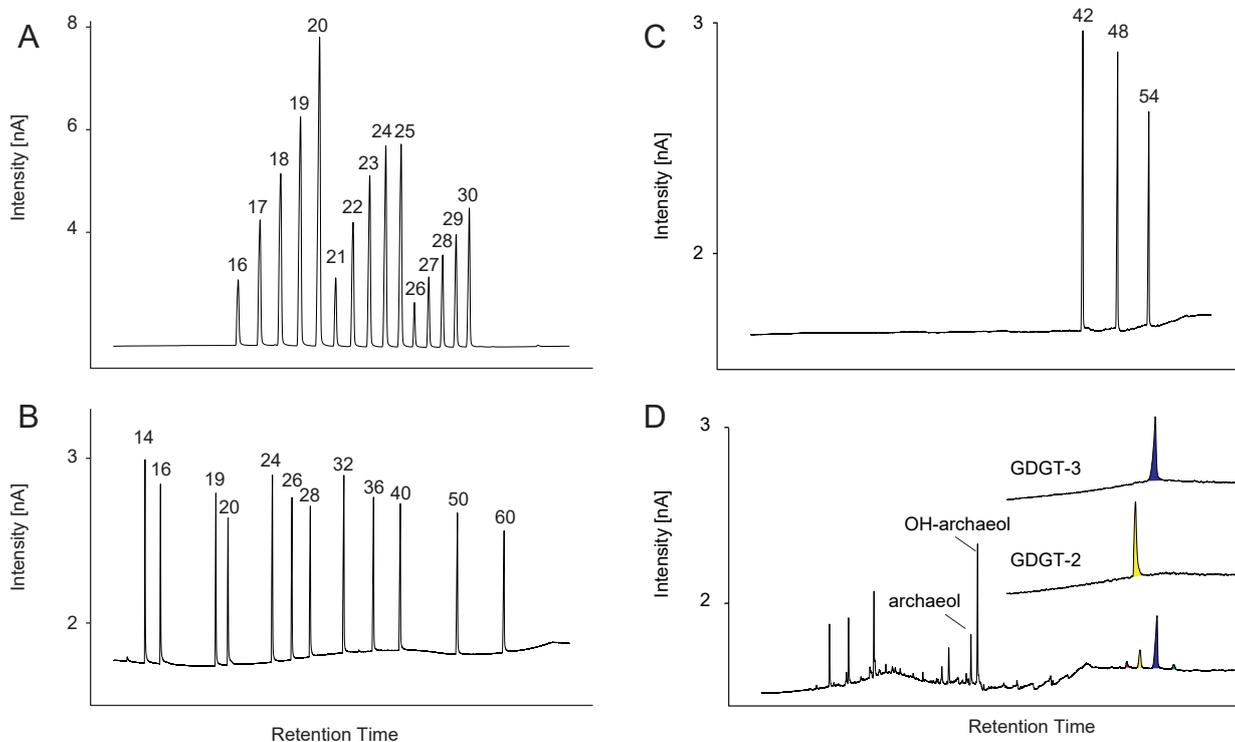


Figure 1. GC-P-IRMS chromatograms under HT conditions, different temperature ramps were applied to the different mixtures. Shown is a mixture of n-alkanes up to n-C30 with known $\delta^2\text{H}$ values (Indiana B3-standard, A), a mixture of long chain n-alkanes up to n-C60 (B), triacylglycerides (C), and a sample from a Black Sea methane seep (D) with GDGT-2 and GDGT-3 standards shown as inserts.

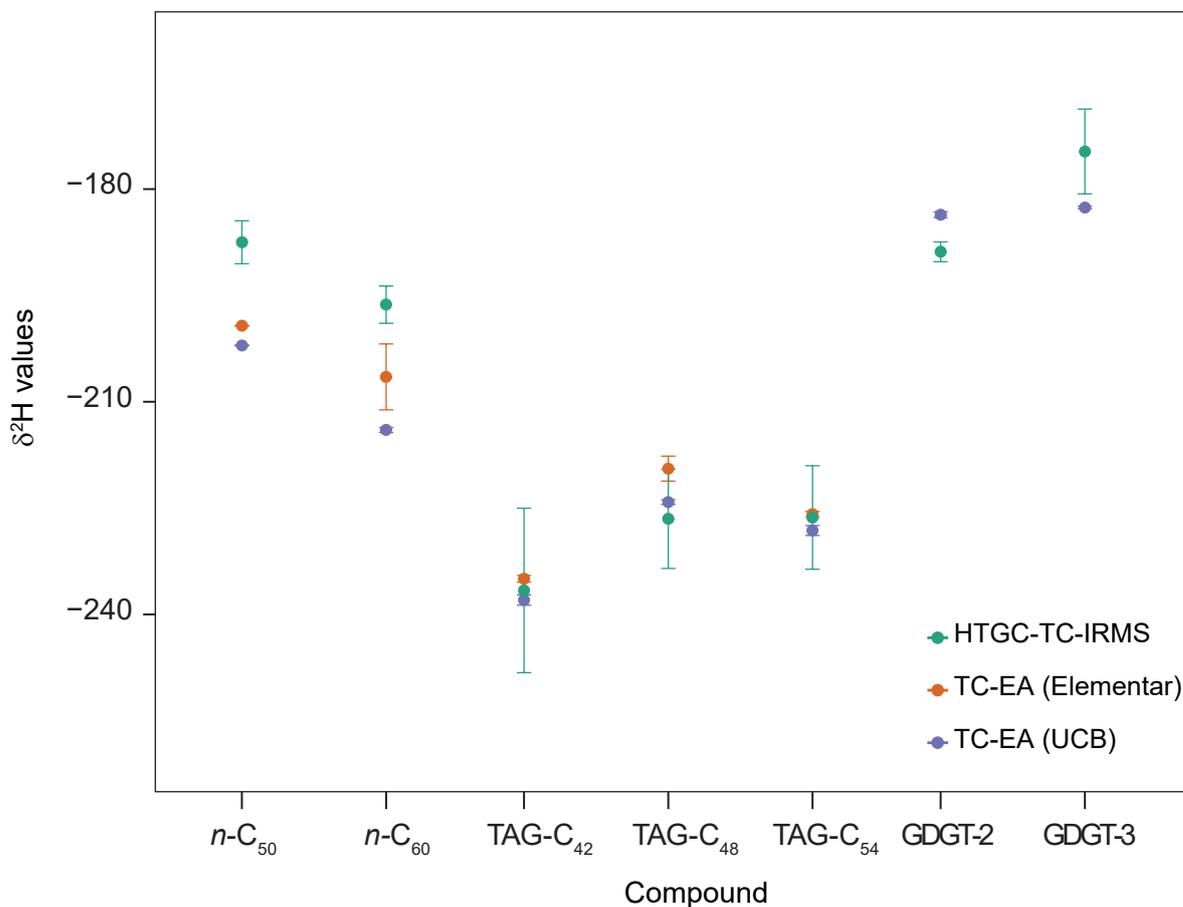


Figure 2. $\delta^2\text{H}$ values of purchased triacylglyceride standards and isolated GDGTs determined by EA-P-IRMS compared with values as determined by HTGC-P-IRMS; values and standard errors are given in Table 1.

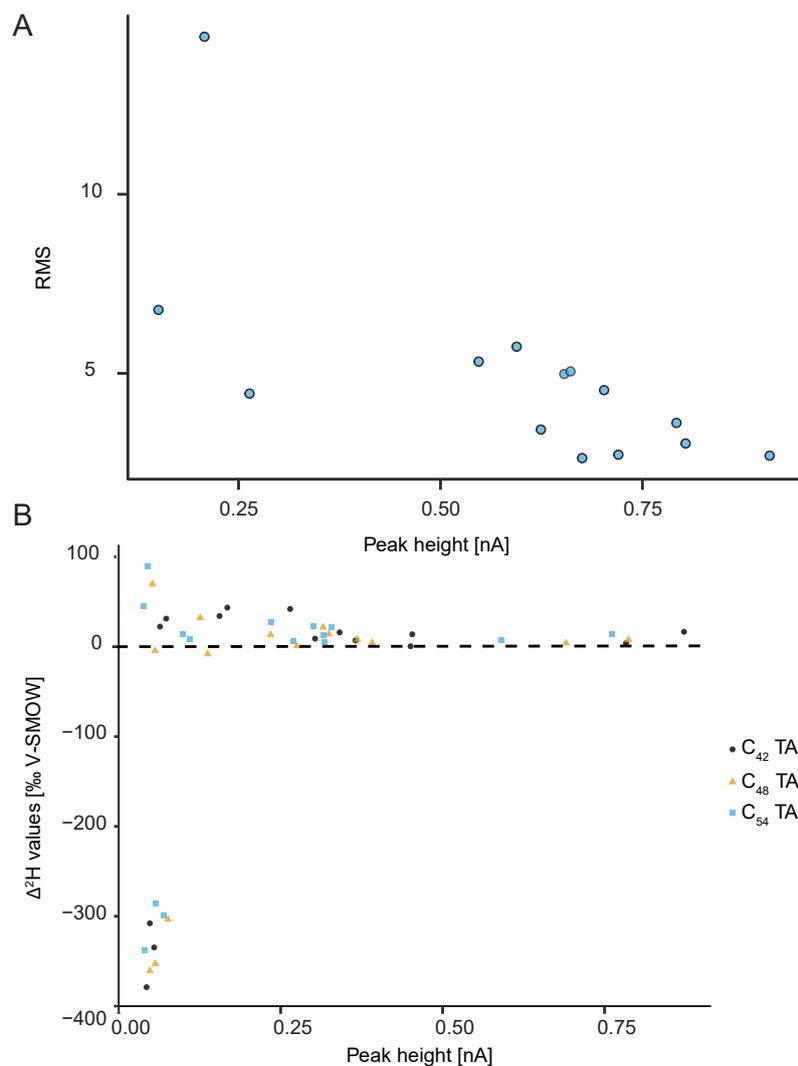


Figure 3. Effect of signal height on $\delta^2\text{H}$ values. A RMSE of the B3 mixture compared to peak heights of the minimum peak height in the mixture. B Difference $\delta^2\text{H}$ values of TAGs determined by HTGC-P-IRMS to values determined by EA-P-IRMS plotted vs peak height.

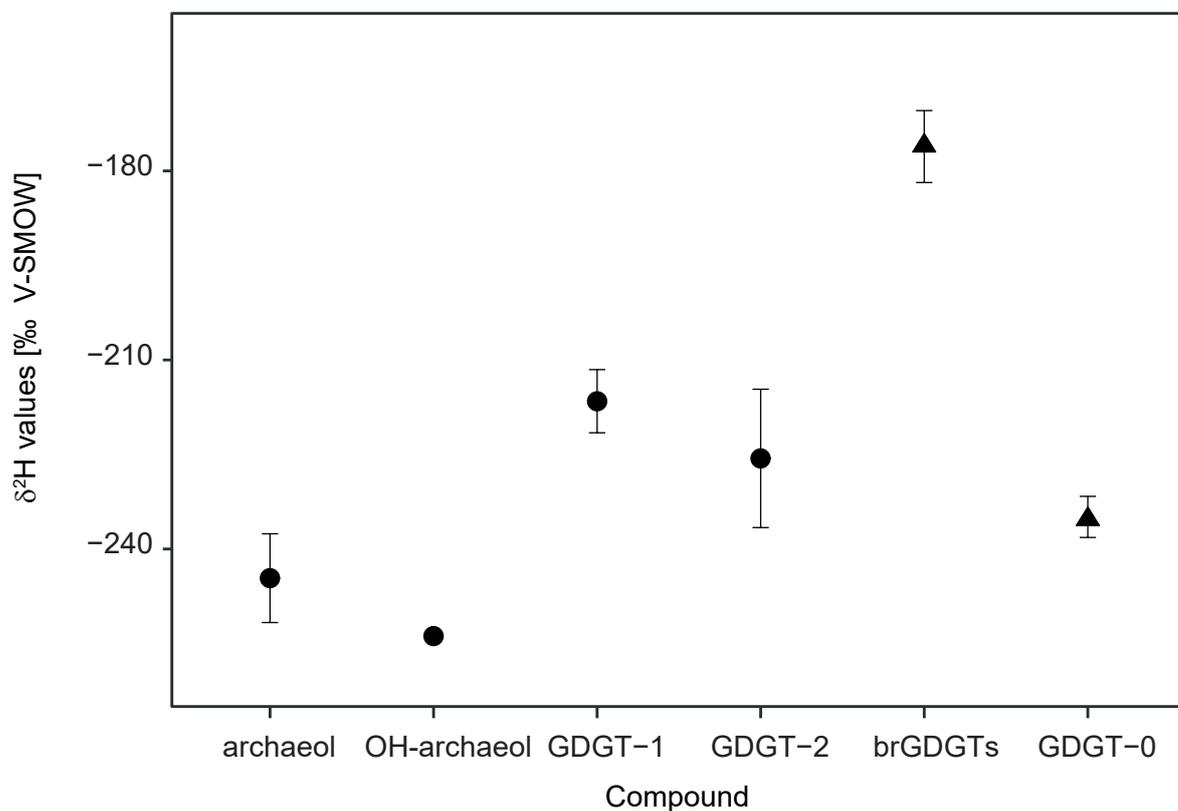


Figure 4. $\delta^2\text{H}$ values of ether lipids determined from environmental samples. brGDGTs and GDGT-0 were extracted from a peat (triangles) and all other compounds derived from a methane seep (circles). Error bars represent standard deviations.

Supplementary to: Determination of the $\delta^2\text{H}$ values of high molecular weight lipids by high temperature GC coupled to isotope ratio mass spectrometry

Sabine K. Lengger^{1,2,*}, Yuki Weber³, Kyle W.R. Taylor⁴, Sebastian H. Kopf⁵, Robert Berstan⁴, Ian D. Bull¹, Jan-Peter Mayser¹, William D. Leavitt⁶, Jerome Blewett¹, Ann Pearson³ and Richard D. Pancost^{1,7}

1 Organic Geochemistry Unit, School of Chemistry, University of Bristol, Cantock's Close, Bristol BS81TS, UK

2 Biogeochemistry Research Centre, School of Geography, Earth and Environmental Science, University of Plymouth, Drake Circus, Plymouth PL48AA, UK

3 Department of Earth and Planetary Sciences, Harvard University, 20 Oxford St, Cambridge, MA 02138, USA

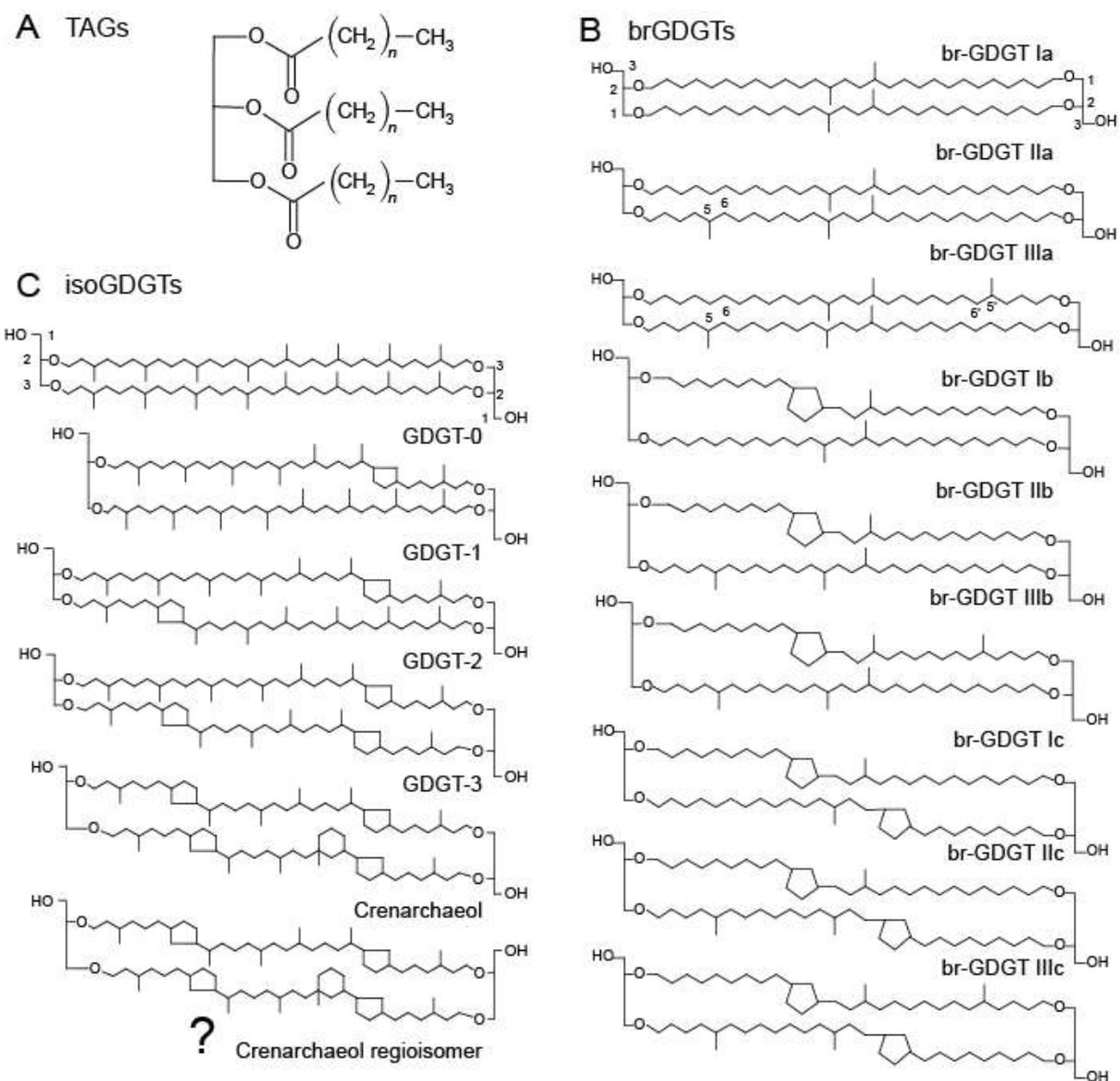
4 Elementar UK Ltd., Earl Road, Cheadle Hulme, Stockport, SK8 6PT, UK

5 Department of Geological Sciences, University of Colorado Boulder, Boulder, CO, USA

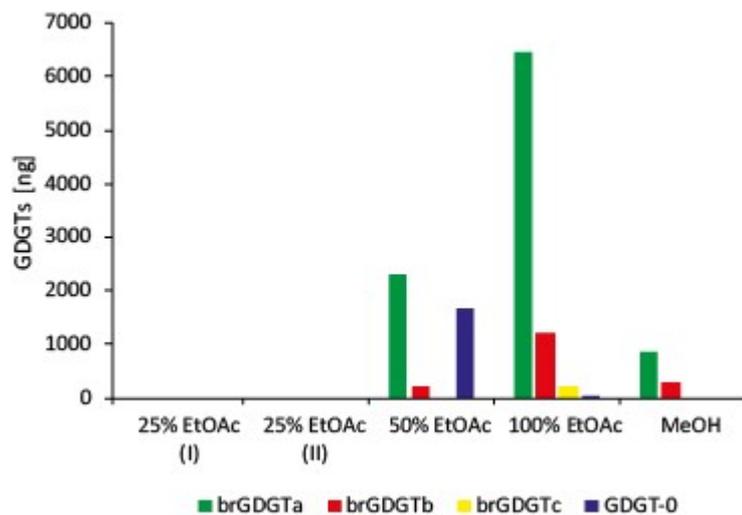
6 Department of Earth Science, Department of Chemistry, Department of Biological Sciences, Dartmouth College, Hanover, NH, USA

7 School of Earth Sciences and Cabot Institute for the Environment, University of Bristol, Queens Road, Bristol BS81RL, UK

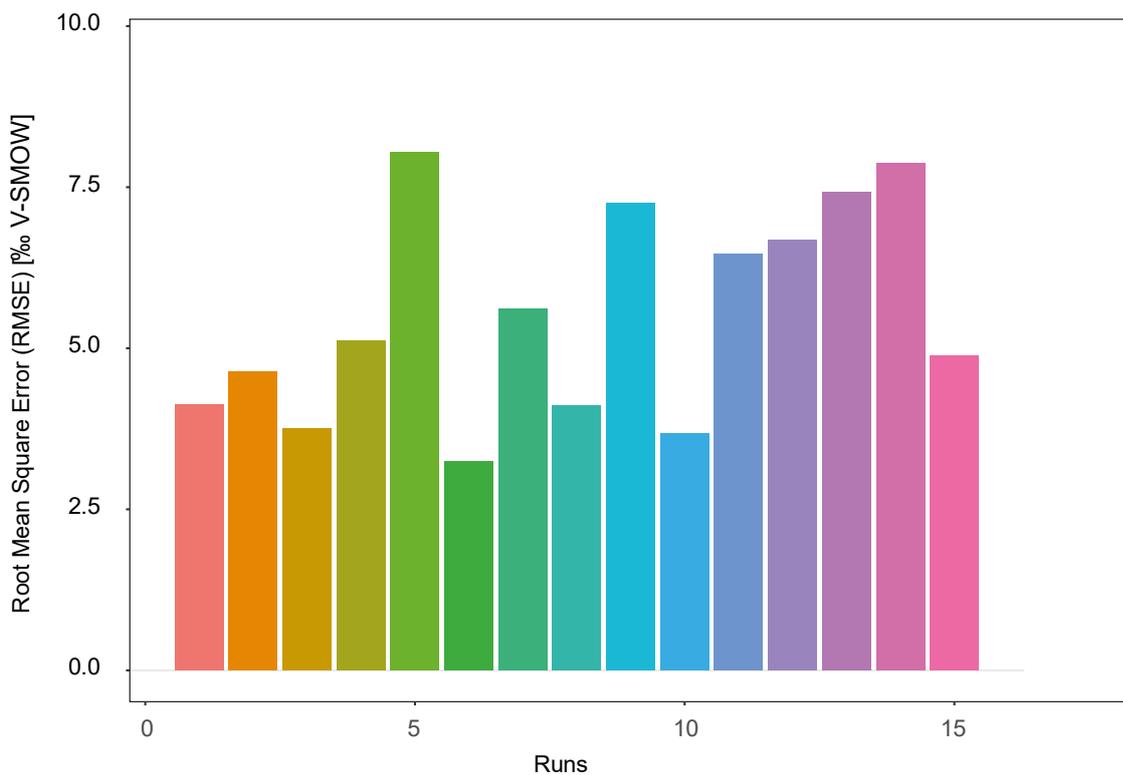
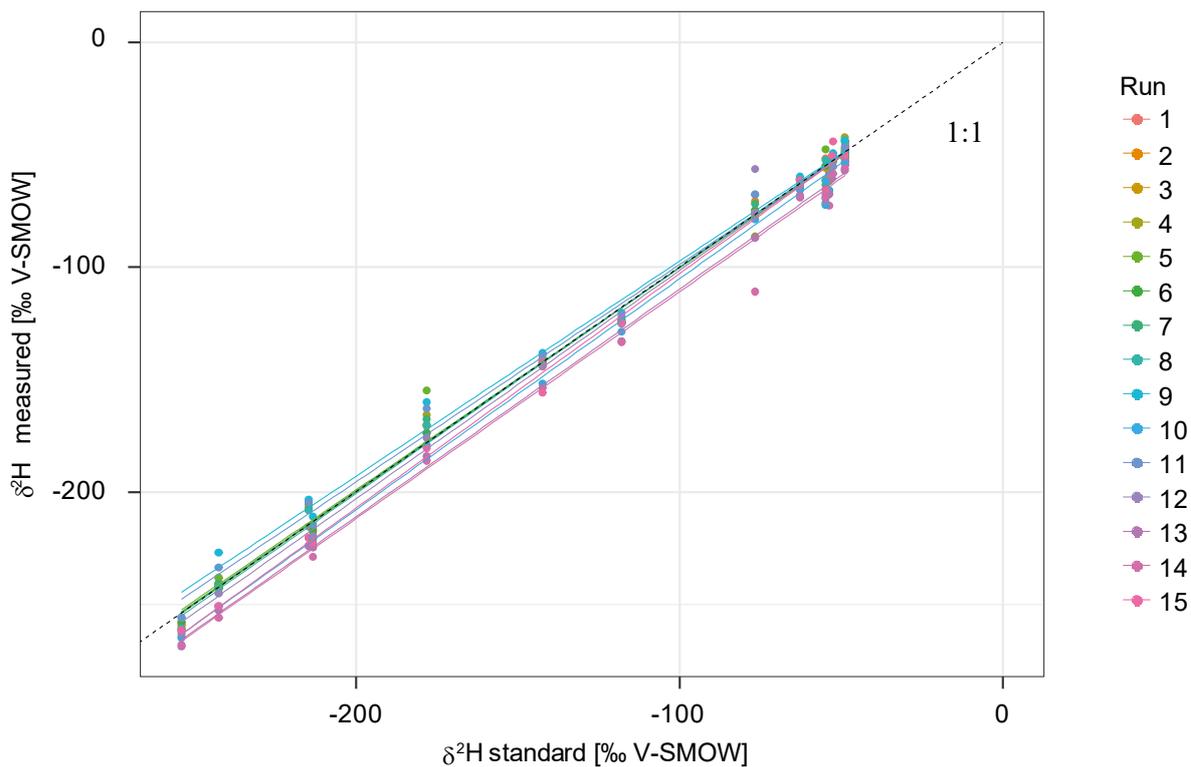
* corresponding author: sabine.lengger@plymouth.ac.uk



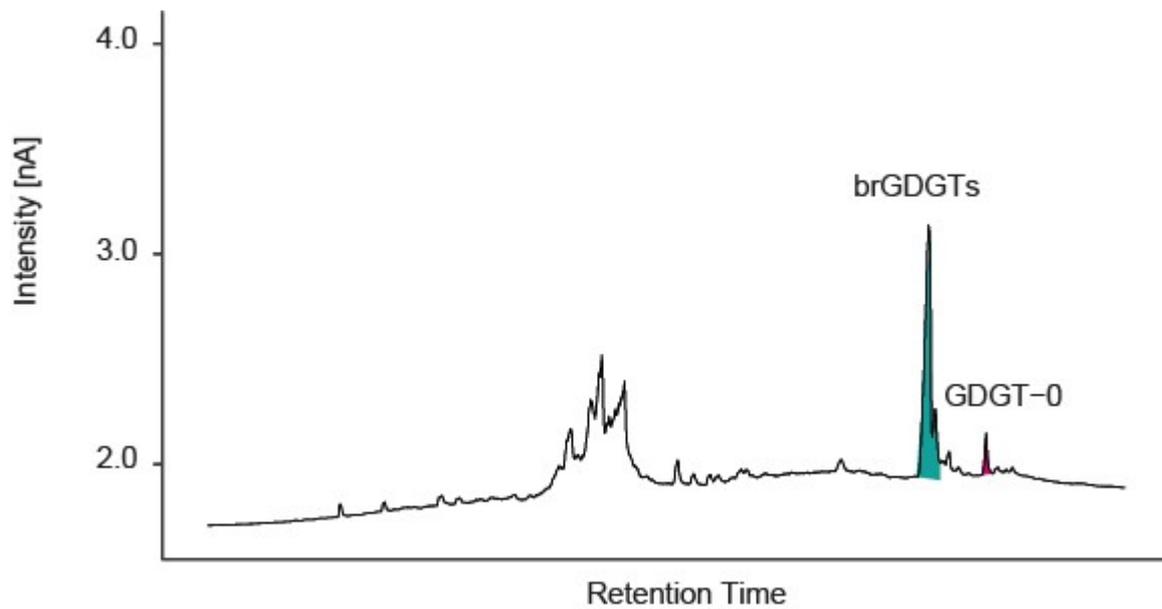
Supplementary Figure 1. Structures of analysed compounds. $n = 14$ (Trimyristate, C_{42}); 16 (Tripalmitate, C_{48}); 18 (Tristearate, C_{54}).



Supplementary Figure 2. GDGTs in column chromatography fractions of sample from a Cors Caron peat (depth 100-200 mm below water table) as determined by HTGC-FID and comparison to triacylglyceride standards as described by Lengger et al. (2018), with a, b and c denoting acyclic, mono-, and bi-cyclic brGDGTs, respectively.



Supplementary Figure 3. Performance monitoring results of *n*-alkane mixture B3 and Root Mean Square Error (RMSE) values of linearisations used for the measurements. Linear regression results and RMSE are shown in table 1.



Supplementary Figure 4. Chromatogram of a combined GDGT-fraction from a Welsh peat analysed by HTGC-P-IRMS.

Table S1. Linearisation coefficients and RMSE of B3 standard runs.

Run nr	Intercept [‰ V-SMOW]	Coefficient	RMSE [‰ V-SMOW]
1	-18.3	0.9615	4.1
2	-16.4	0.9629	4.6
3	-20.2	0.9759	3.8
4	-19.4	0.9152	5.1
5	-21.0	0.9120	8.0
6	-18.6	0.9924	3.2
7	-20.1	0.9597	5.6
8	-19.0	0.9703	4.1
9	-21.1	0.9225	7.3
10	-21.9	0.9905	3.7
11	-21.7	0.9319	6.5
12	-19.4	0.9799	6.7
13	-27.7	0.9762	3.0
14	-29.2	0.9730	7.9
15	-18.2	1.0048	4.9